Interannual variability of the western boundary current of the Bay of Bengal: 25 years of repeated XBT sections

Sherin V Raju, Fabien Durand, Chaitanya A V S, Gopalkrishna V V

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presented by : Dr.V.V. GopalKrishna National Institute of Oceanography Goa , India, gopal@nio.org

**Background of the Study area (Bay of Bengal)** 



- Semi enclosed tropical basin, restricting northward transport of Heat & Mass.
- Evaporation Précipitation is always ve
- Strong haline stratification, shallow mixed layer,

4<sup>o</sup> - 5<sup>o</sup>C Temp. Invensions

Devastating Tropical Cyclones



- Seasonal riversal of monsoonal winds, July (Southwesterlies) & January (Northeasterlies)
- Western BoB is prone to strong mesoscale eddies; influencing the coastal currents.

Background of the study area (Bay of Benga

- East India Coastal Current (EICC). Poleward (Feb Sep) & Equatorward (Oct Jan)
- EICC participates in the Exchange of Water Masses between basins.



Summer

Winter

(Schott & McCreary, 2001)

# **XBT** Observations

![](_page_4_Picture_1.jpeg)

![](_page_4_Picture_2.jpeg)

- T-7 XBT probes
- 1990 continuing (25 years)
- Monthly voyages
- SSS Samples: at all XBT Stations

![](_page_4_Figure_7.jpeg)

XBT data density (25 years

25'

20'

15'

10'

5°

0° L 70° INDIA

Chen

75°

80"

85"

BoB

90"

15

Port Blair

95"

100

![](_page_5_Figure_1.jpeg)

## Chennai – Port Blair Transect : 1990 - continuing

Computational strategy of geostrophic cu

- Carried out running averages for 1<sup>o</sup>x1<sup>o</sup> bin for each XBT transect.
- Monthly T-S relationship is developed from NIOA climatology.
- Using this T-S relation, salinity values are synthesized for each XBT (t) values.

![](_page_6_Figure_4.jpeg)

![](_page_6_Figure_5.jpeg)

![](_page_6_Figure_6.jpeg)

Computational strategy of geostrophic cu

Density profiles are computed for each XBT (t) & Synthesized salinity.

Crosstrack geostrophic currents are computed w.r.t 700m reference level

$$V = \frac{g}{f\rho} \frac{\partial}{\partial x} \int_{-h}^{0} \rho(z) \, dz$$

• A Hanning filter is applied to remove small scale fluctuations

![](_page_7_Figure_5.jpeg)

![](_page_7_Figure_6.jpeg)

![](_page_8_Figure_0.jpeg)

**Boundary current is flowing northward during Summer (Dark shade** 

**Eequatorward during Winter (Blue shade)** 

![](_page_8_Figure_3.jpeg)

Seasonal climatology of geostrophic currents

![](_page_9_Figure_1.jpeg)

![](_page_10_Figure_0.jpeg)

![](_page_11_Figure_0.jpeg)

Volume transport shows seasonal reversal (southward) during post monsoon. Varies between 5Sv northward & 3Sv southward
Time window of equatorward flow which transports low salinity waters is small.

## What drives this Inter- annual Variability of the Western Boundary Current Transports ??

# Whether IOD events influence the currents in the Bay of Bengal **??**

![](_page_12_Figure_2.jpeg)

#### Results

## The Dipole Mode Index (DMI)

Saji *et al.* (1999) define DMI as the difference between the SST anomalies in the western equatorial Indian Ocean (50°E to 70°E, 10°S to 10°N, Western basin) and the southeastern tropical Indian Ocean (90°E to 110°E, 10°S to the Equator, Eastern basin).

Positive Dipole Mode

![](_page_13_Picture_4.jpeg)

As per the above, definition the following years are considered as<sup>3.0</sup> +ve IOD & -ve IOD years; 1991, 94, 97, 2006,11,12,15. <sup>2.0</sup> 1992, 96, 98, 2010.

Circles represents IOD years,  $\frac{\circ}{3}$ Where we have XBT data.

![](_page_13_Figure_7.jpeg)

![](_page_14_Figure_0.jpeg)

![](_page_14_Figure_1.jpeg)

![](_page_14_Figure_2.jpeg)

What can 25 years of XBT repeated sections tell us ?...

#### Results

#### Composites of current anomalies for IOD & Non – IOD years: Sep, Oct, Nov

![](_page_15_Figure_2.jpeg)

•The anomaly composites shows clear signature of the influence of IOD events in the WBC.

 +ve IOD composite shows northward anomaly & -ve IOD southward anomaly.

•For normal years the anomaly is close to zero.

• The anomalous currents are narrow & at least 700m deep.

![](_page_15_Figure_7.jpeg)

#### **Composites of Transport** for IOD Years

![](_page_16_Figure_1.jpeg)

•The influence of IOD in the WBC is manifested by anomalous poleward transport (5.21 Sv, +ve IOD) & anomalous equatorward transport (6.56 Sv, -ve IOD ).

• The transport anomaly for +ve IOD and –ve IOD are distinct from each other after accounting for the standard deviations.

### Results

![](_page_17_Figure_1.jpeg)

Low correlation (0.24) between DMI time Series & WBC Variability indicates that, although Western Boundary Current transports showed distinct signature of + ve IOD & - ve IOD events, it appears that most of the Year-to-Year variability is not related to IOD events.

Other factors contributing to the Variability looks to be;

![](_page_18_Figure_2.jpeg)

- Mesoscale Turbulence
- ➢Instabilites of the WBC
- Local Winds

The western boundary of the BoB: a key region of the tropical climate variability.

- 25 years of repeat XBT sections: an unprecedented coverage in a data-sparse region.
- The variability of the WBC is dominated by remote forcing from equatorial ocean, during IOD years.
- This manifests as narrow, deep (700m) anomalous transport
- WBC Location: off the shelf break slope region.
- We must have Pluri Annual monitoring of WBC transport

![](_page_19_Figure_7.jpeg)

Aside from remote forcing, local forcing (turbulence / local winds) apparently prominent in WBC modulation.

![](_page_20_Picture_0.jpeg)